

Normal modes

Standing waves 1-D

We start from the wave equation

$$\frac{\partial^2 u}{\partial t^2} = c^2 \frac{\partial^2 u}{\partial x^2}$$

And use separation of variables
resulting in

$$u(x, t) = X(x)T(t)$$

$$\frac{1}{T} \frac{d^2 T}{dt^2} = \frac{c^2}{X} \frac{d^2 X}{dx^2} = -\omega^2$$

We solve for X and T separately, giving

$$T(t) = \begin{bmatrix} \cos(\omega t) \\ \sin(\omega t) \end{bmatrix}$$

$$X(x) = \begin{bmatrix} \cos(\omega x/c) \\ \sin(\omega x/c) \end{bmatrix}$$

Standing waves 1-D

The general solution is

$$u(x, t) = C_1 \cos(\omega t) \cos(\omega x/c) + C_2 \cos(\omega t) \sin(\omega x/c) + C_3 \sin(\omega t) \cos(\omega x/c) + C_4 \sin(\omega t) \sin(\omega x/c)$$

We use the boundary conditions

$$(1) \quad u(0, t) = 0 \rightarrow C_1 = C_3 = 0$$

$$(2) \quad u(L, t) = 0 \rightarrow \sin(\omega L/c) = 0$$

$$\frac{\omega L}{c} = (n + 1)\pi$$

$$\omega_n = \frac{(n + 1)\pi c}{L}$$

**Discrete
eigenfrequencies**

Standing waves 1-D

$$\omega_n = \frac{(n+1)\pi c}{L}$$

**Discrete
eigenfrequencies**

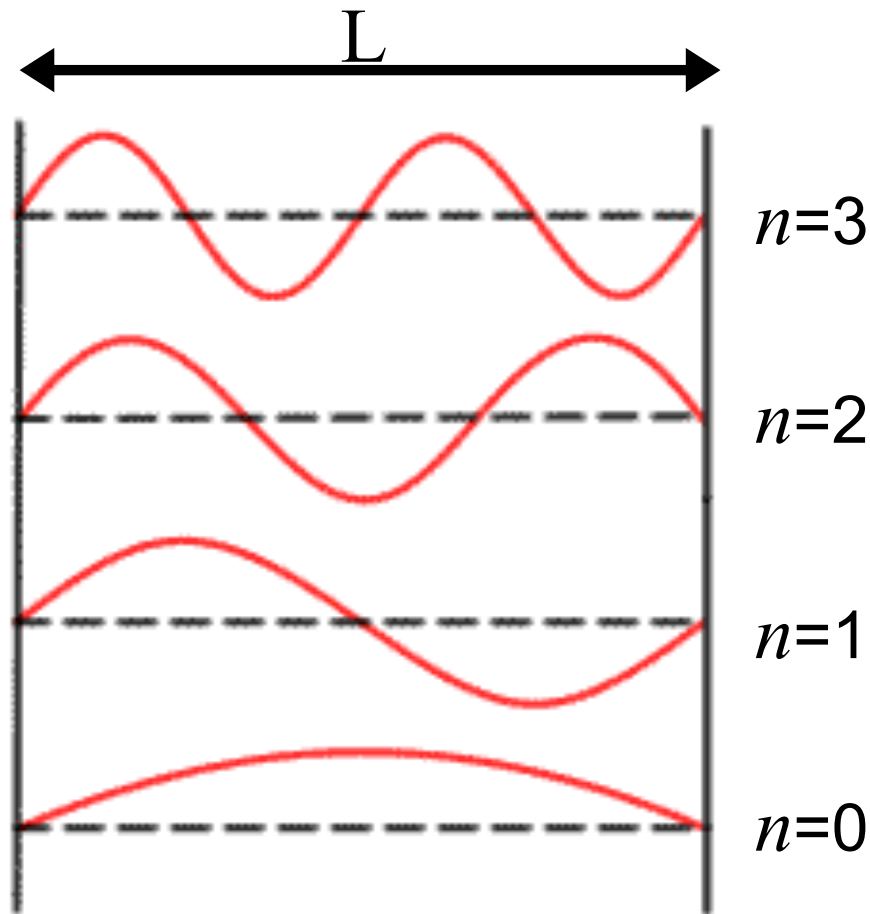
$$\sin\left(\frac{\omega_n x}{c}\right)$$

**Eigenfunctions or
normal modes**

Solutions are of the form

$$u(x, t) = \sum_{n=0}^{\infty} [A_n \cos(\omega_n t) + B_n \sin(\omega_n t)] \sin\left(\frac{\omega_n x}{c}\right)$$

Standing waves 1-D



Standing waves on a string

Frequency of standing wave on a string:

$$\omega_n = \frac{(n+1)\pi c}{L}$$

where c = wave velocity

Fourier spectrum will have spikes at ω_n

Normal mode summation

The amplitudes A_n depend on the source

$$A_n = \sin(n\pi x_s/L)F(\omega_n)$$

So, the normal mode summation is

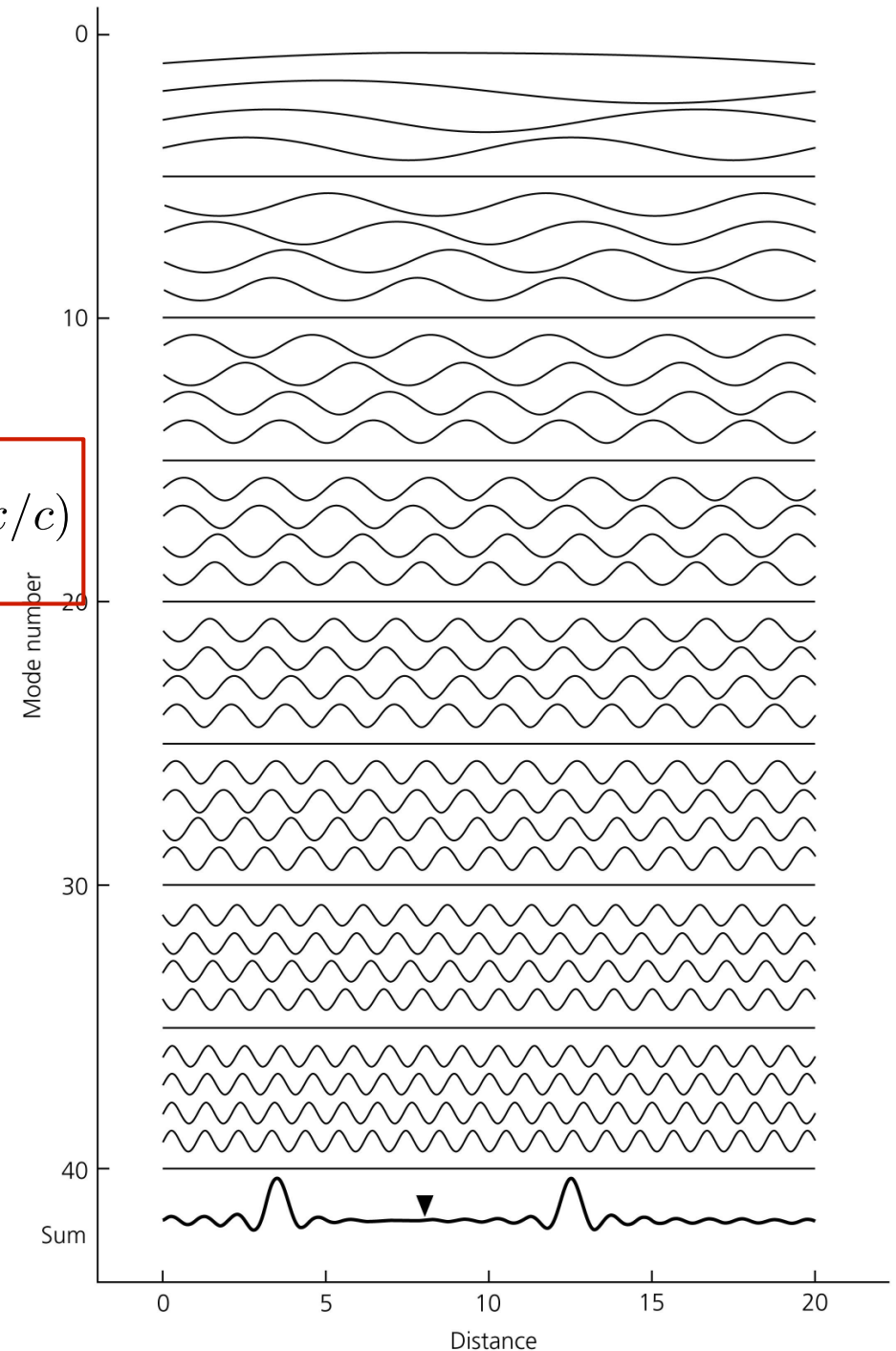
$$u(x, t) = \sum_{n=0}^{\infty} \sin(n\pi x_s/L) F(\omega_n) \cos(\omega_n t) \sin(\omega_n x/c)$$

with the source, at $x_s=8$, described by

$$F(\omega_n) = \exp[-(\omega_n \tau)^2/4]$$

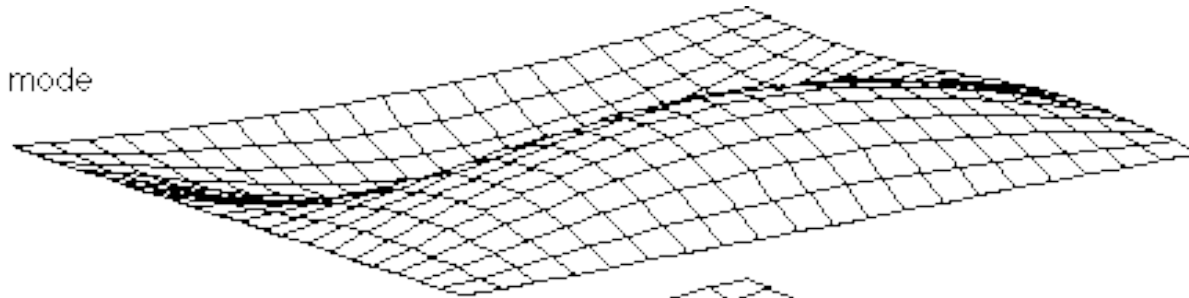
for $\tau = 0.2$

Figure 2.2-8: Waves on a string as a summation of modes.

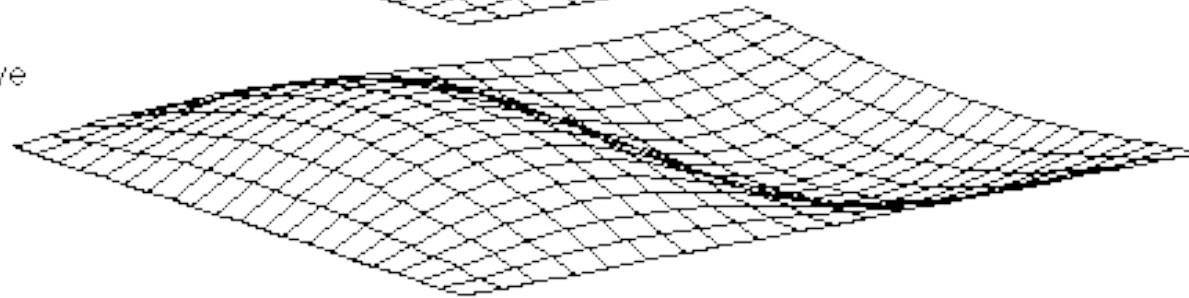


Standing waves 2-D

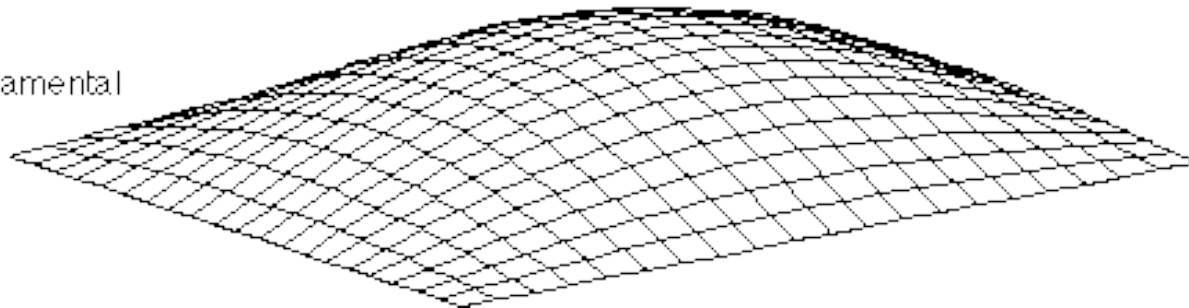
Third mode



Octave

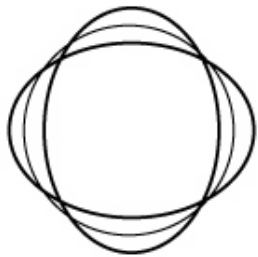


Fundamental

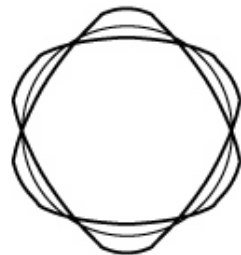


Standing waves 3-D

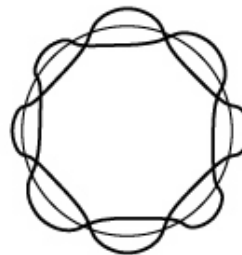
Surface patterns



${}_0S_2$



${}_0S_3$



${}_0S_4$

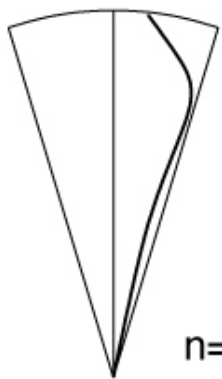
Spherical harmonics

$$Y_l^m(\theta, \phi)$$

Angular order l
 $l=0,1,2,\dots$

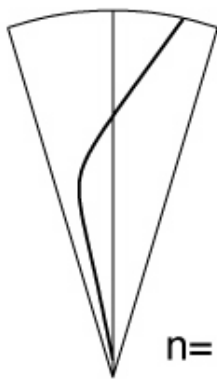
Azimuthal order m
 $m=-l,-l+1,1,\dots,0,\dots,l$

Radial patterns



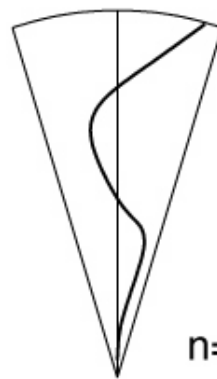
$n=0$

Fundamental



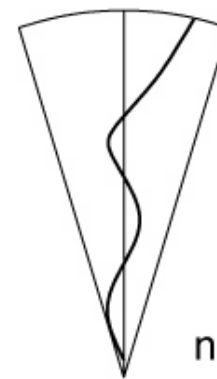
$n=1$

First Overtone



$n=2$

Second Overtone



$n=3$

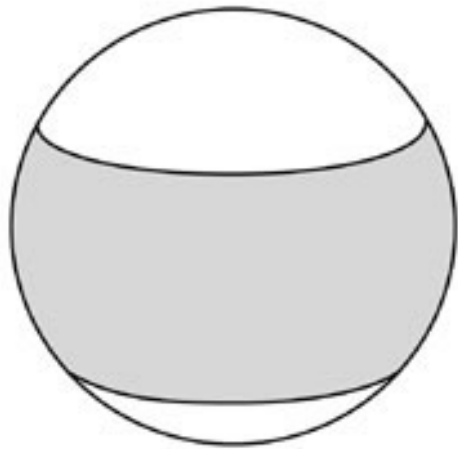
Third Overtone

Spherical Bessel function

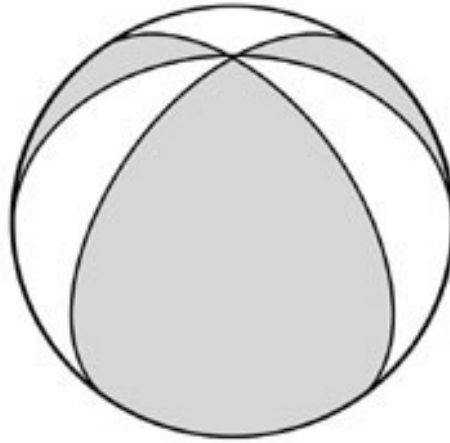
$$j_n(kr)$$

Radial order n
 $n=0,1,2,\dots$

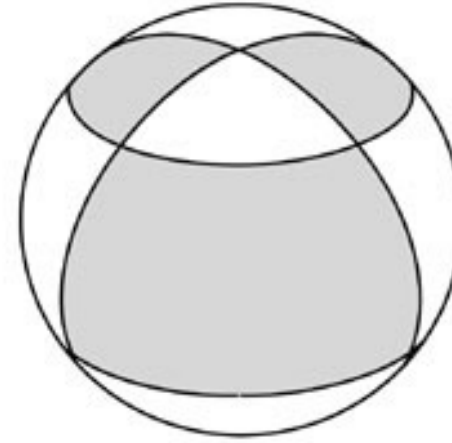
Spherical harmonics



Y_2^0



$\text{Re}(Y_3^3)$



$\text{Re}(Y_4^2)$

$$Y_l^m(\theta, \phi) = (-1)^m \left[\left(\frac{2l+1}{4\pi} \right) \frac{(l-m)!}{(l+m)!} \right]^{1/2} P_l^m(\cos \theta) e^{im\phi}$$

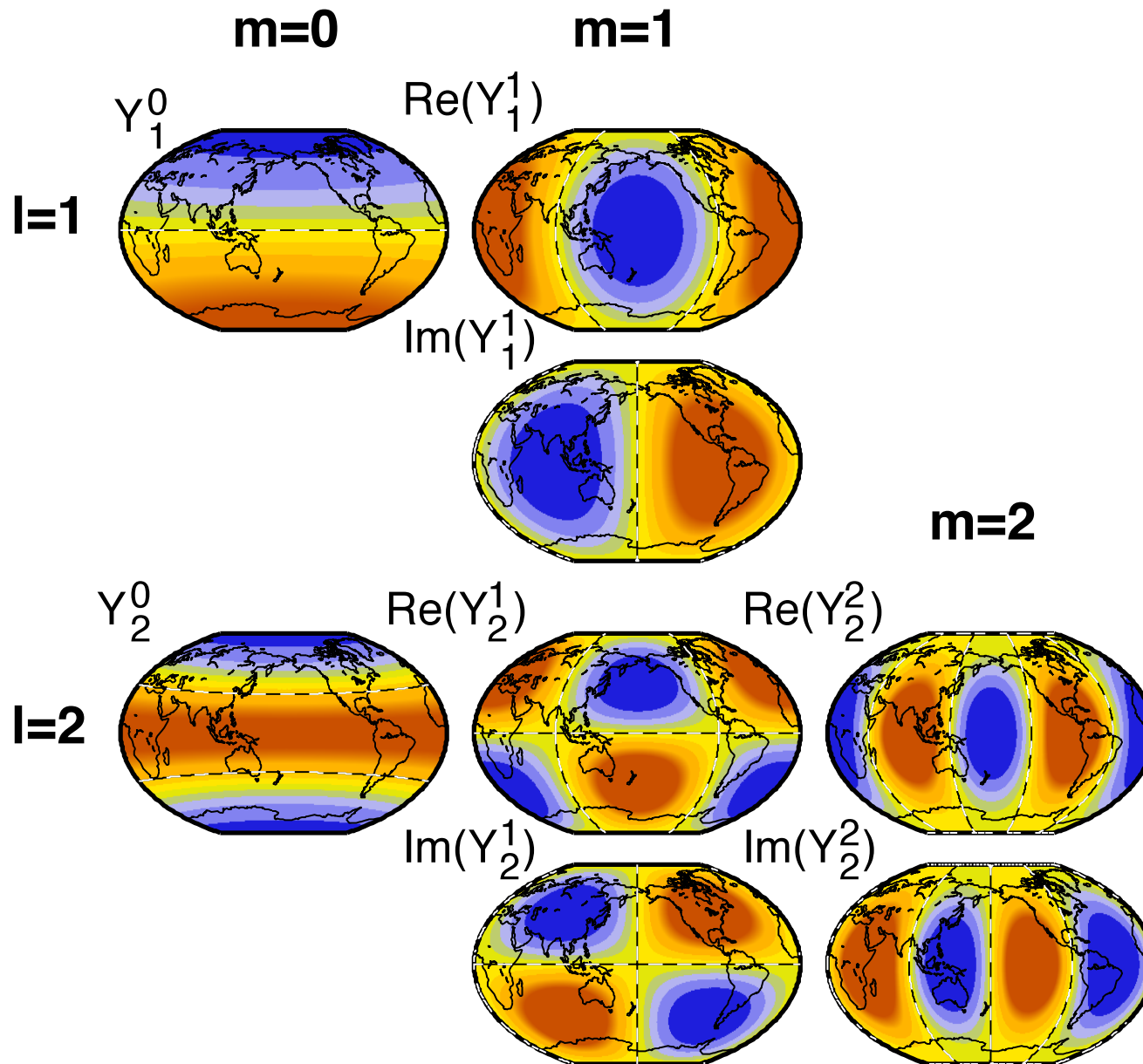
(where $P_l^m(\cos \theta)$ is associated Legendre function)

The **angular order** l gives the number of nodal lines on the surface

If the **azimuthal order** m is zero, the nodal lines are small circles about the pole. These are called zonal harmonics and do not depend on ϕ .

For a given angular order l , m has $2l+1$ values, leading to $2l+1$ different singlets (or eigenfunctions)

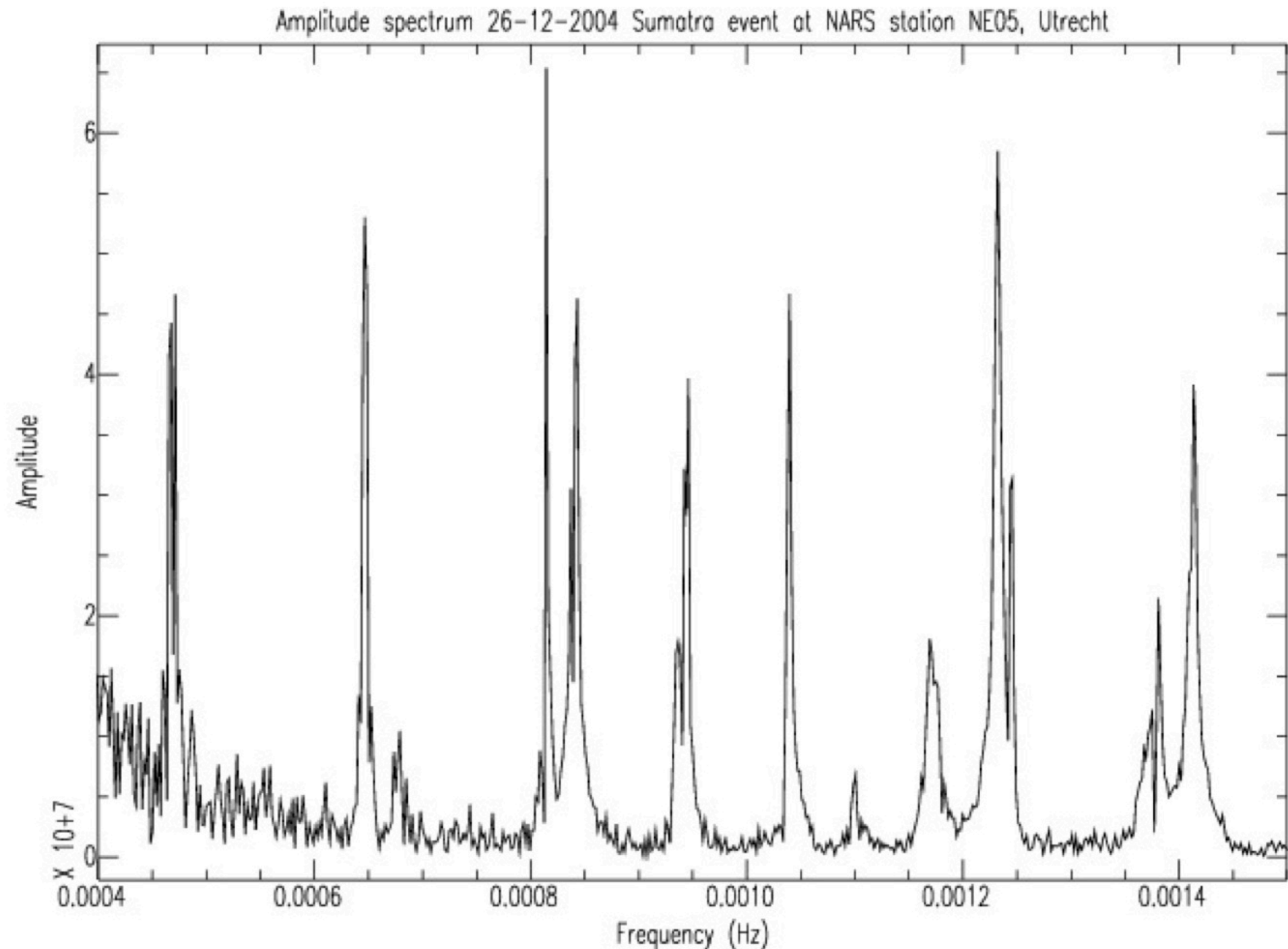
Spherical harmonics



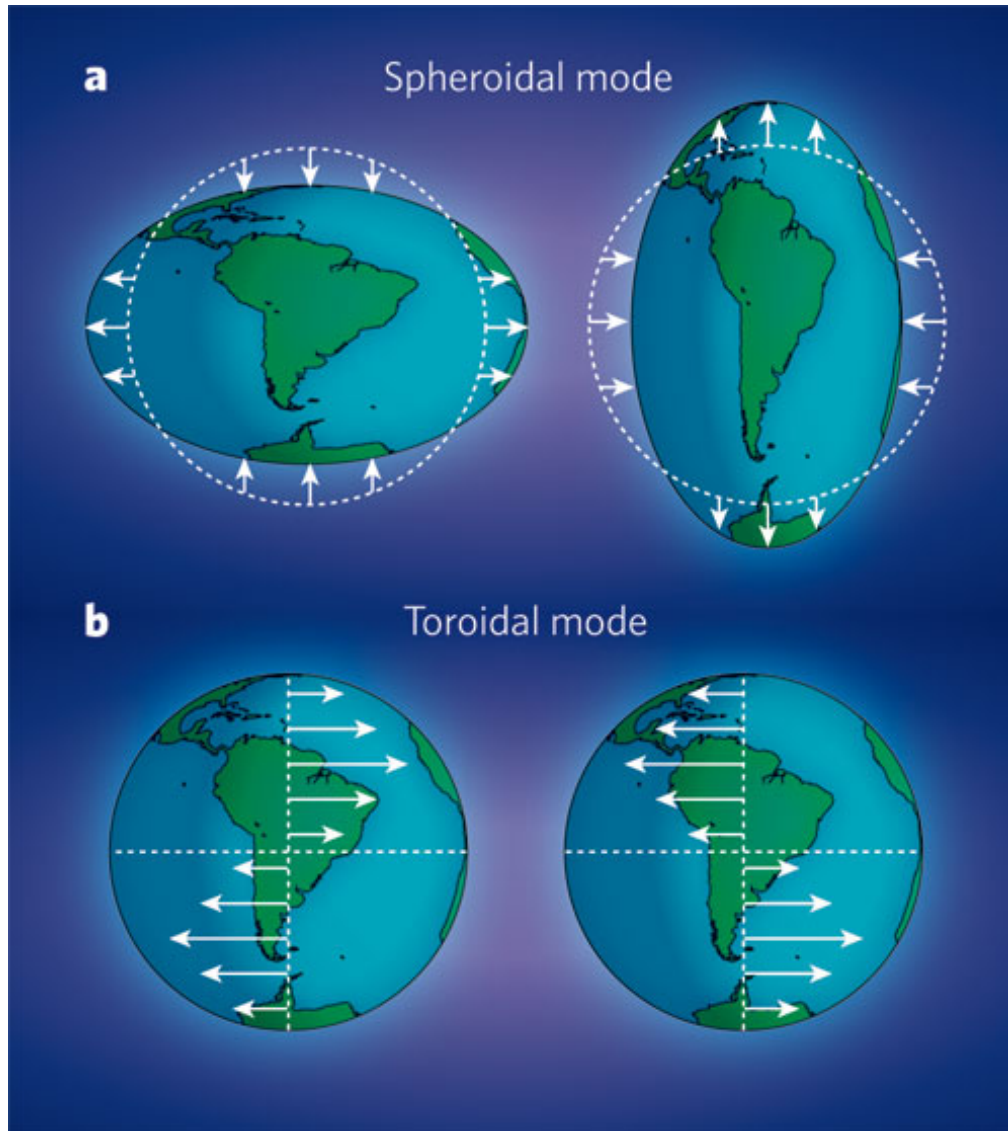
Sumatra earthquake 2004, M 9.1



Sumatra earthquake recorded in Utrecht

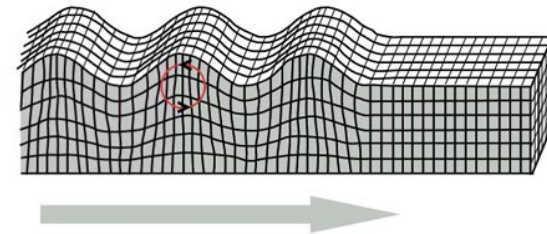


Displacement direction



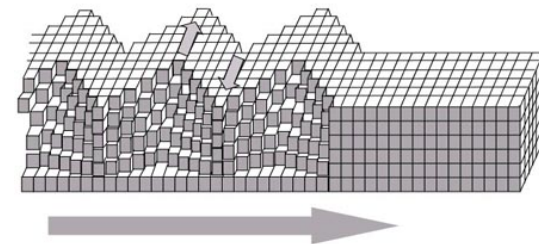
Spheroidal modes

- P-SV motion
- Similar to Rayleigh



Toroidal modes

- SH motion
- Similar to Love



Toroidal modes

For ${}_nT_l^m$:

n =radial order, l =angular order, m =azimuthal order

The $2l+1$ modes are different azimuthal orders $l=-m, -m+1, \dots, 0, \dots, m$ are called singlets, and the group of singlets is called a multiplet.

If the Earth were perfectly spherically symmetric and non-rotating, all singlets in a multiplet would have the same frequency (called degeneracy).

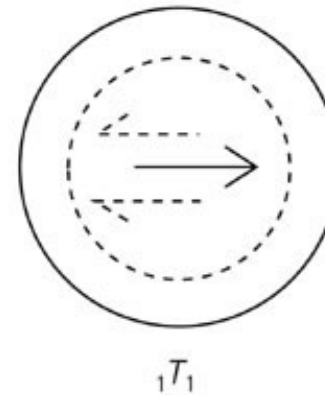
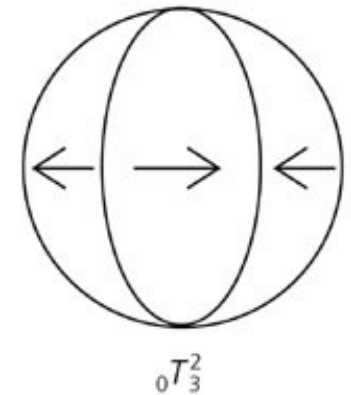
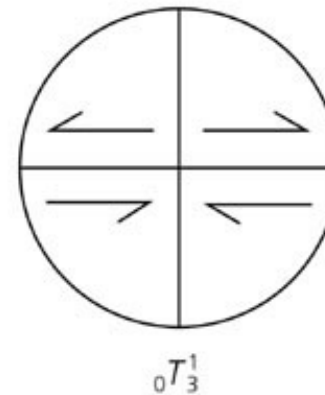
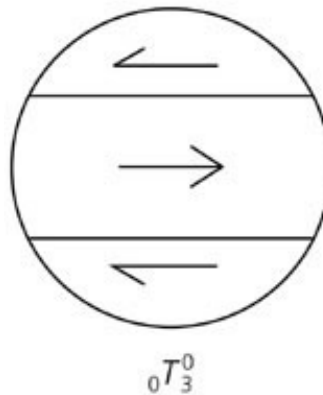
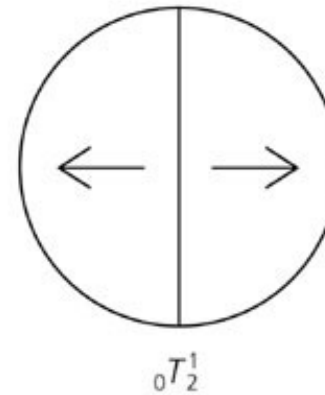
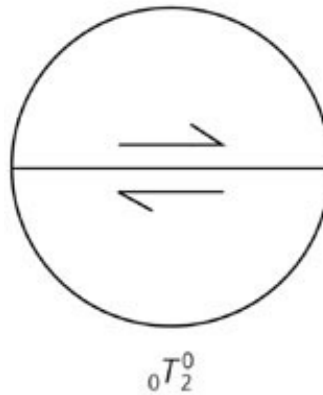
For example, the period of ${}_nT_l^0$ would be the same for ${}_nT_l^1, {}_nT_l^2$ etc. In the real Earth, singlet frequencies vary (called splitting).

The splitting is usually small enough to ignore, so we drop the m superscript and refer to the entire ${}_nT_l^m$ multiplet at ${}_nT_l$ with eigenfrequency ${}_n\omega_l$

Toroidal modes

Toroidal modes with $n=0$ (${}_0T_l$) are called **fundamental modes** (motions at depth in the same direction as at the surface).

Modes with $n>0$ are called **overtones** (motions reverse directions at different depths).



Spheroidal modes

For ${}_nS_l^m$:

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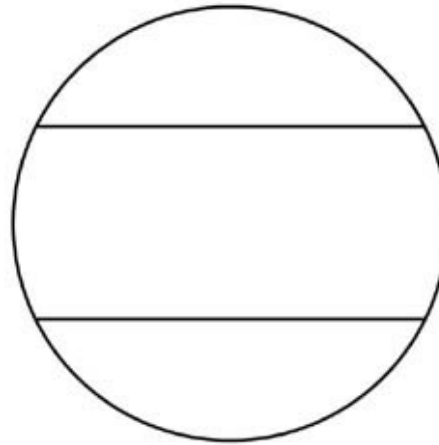
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Spheroidal modes

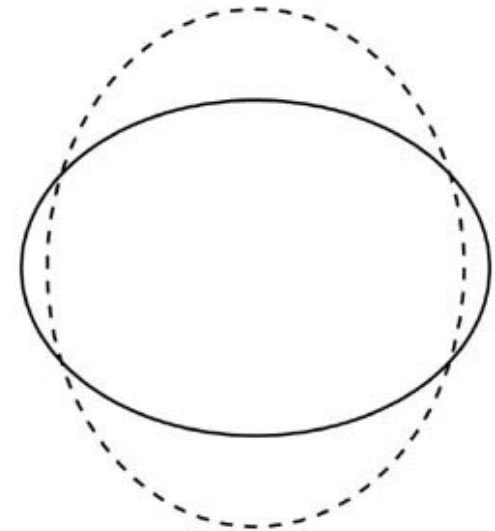
${}_0S_2$ (football mode) is the gravest (lowest frequency) mode, with a period of 3233 seconds, or 54 minutes.

There is no ${}_0S_1$ which would correspond to a lateral translation of the planet.

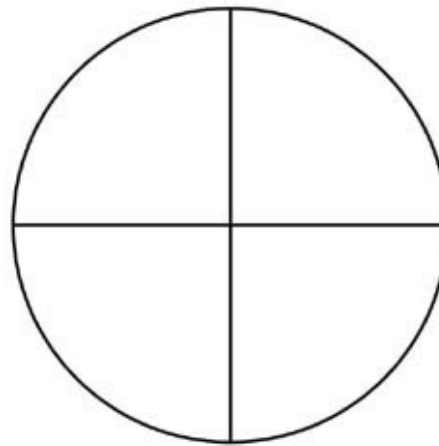
The ${}_1S_1$ Slichter mode due to lateral sloshing of the inner core through the liquid outer core, is not yet observed, but should have a frequency of about 5 ½ hours.



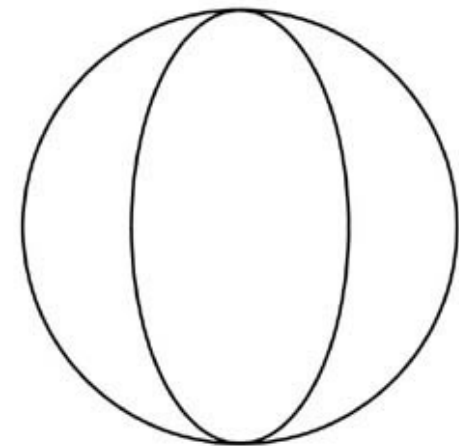
${}_0S_2^0$



${}_0S_2^0$ (motion)

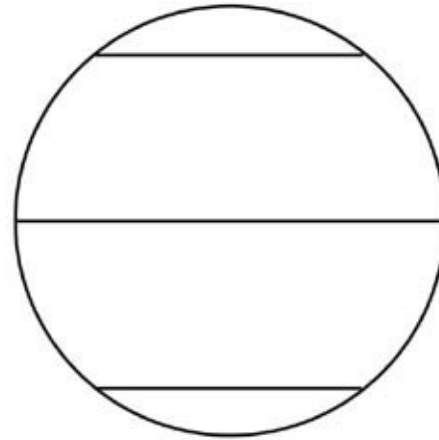


${}_0S_2^1$

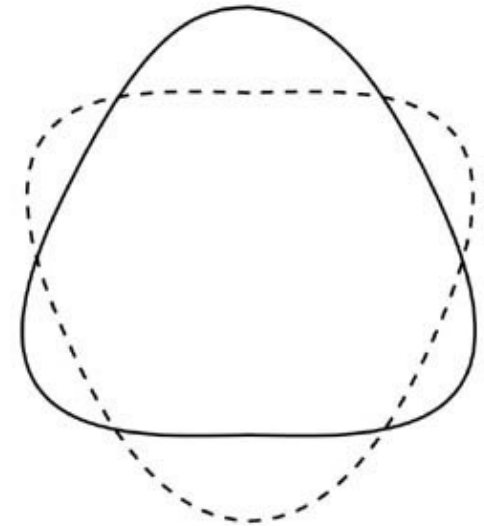


${}_0S_2^2$

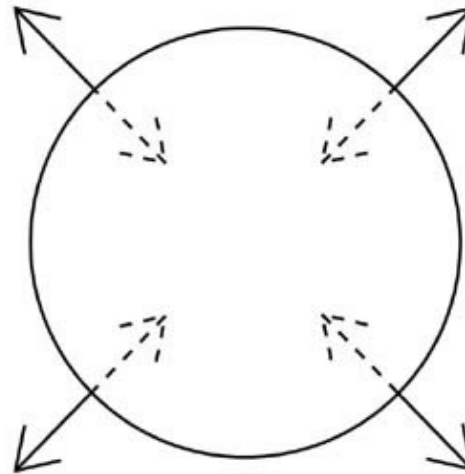
Spheroidal modes



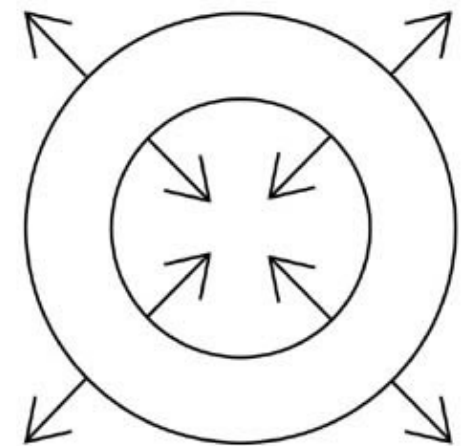
${}_0S_3$



${}_0S_3$ (motion)



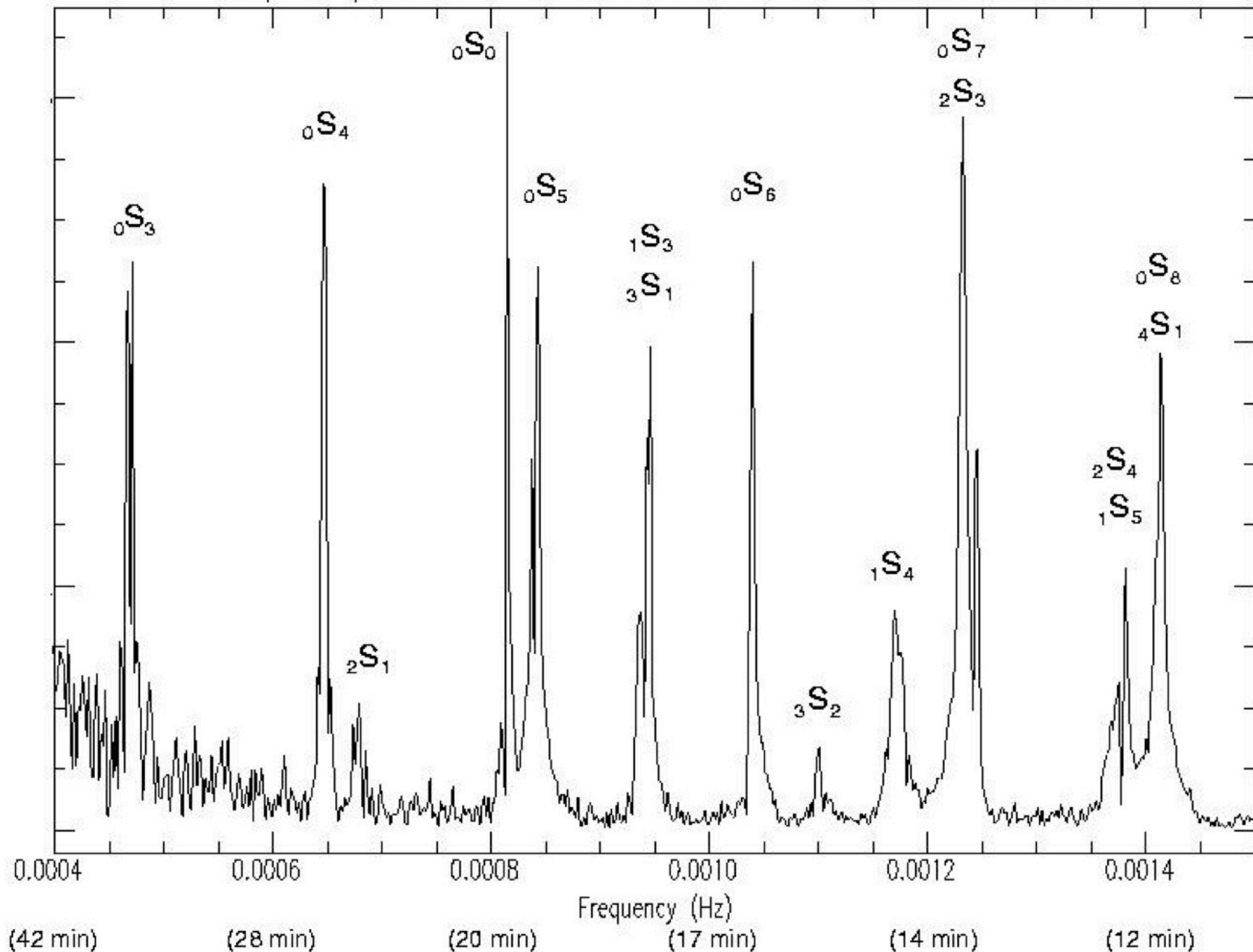
${}_0S_0$



${}_1S_0$

The “breathing” mode ${}_0S_0$ involves radial motions of the entire Earth that alternate between expansion and contraction.

Amplitude spectrum 26-12-2004 Sumatra event at NARS station NE05, Utrecht



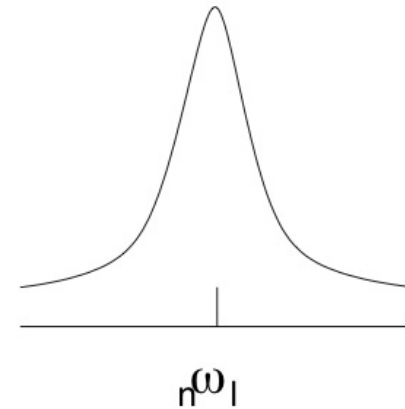
Eigenfrequencies of the earth excited by the Sumatra earthquake and recorded in Utrecht.

SNREI

- * spherical symmetric,
- * non-rotating,
- * isotropic Earth

degeneracy:

$2l+1$ singlets have same frequency



Real Earth

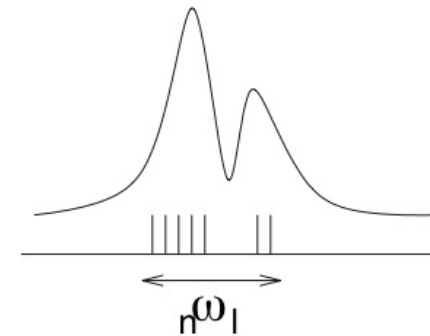
- * Rotation
- * Ellipticity
- * Heterogeneity
- * Anisotropy

degeneracy removed:

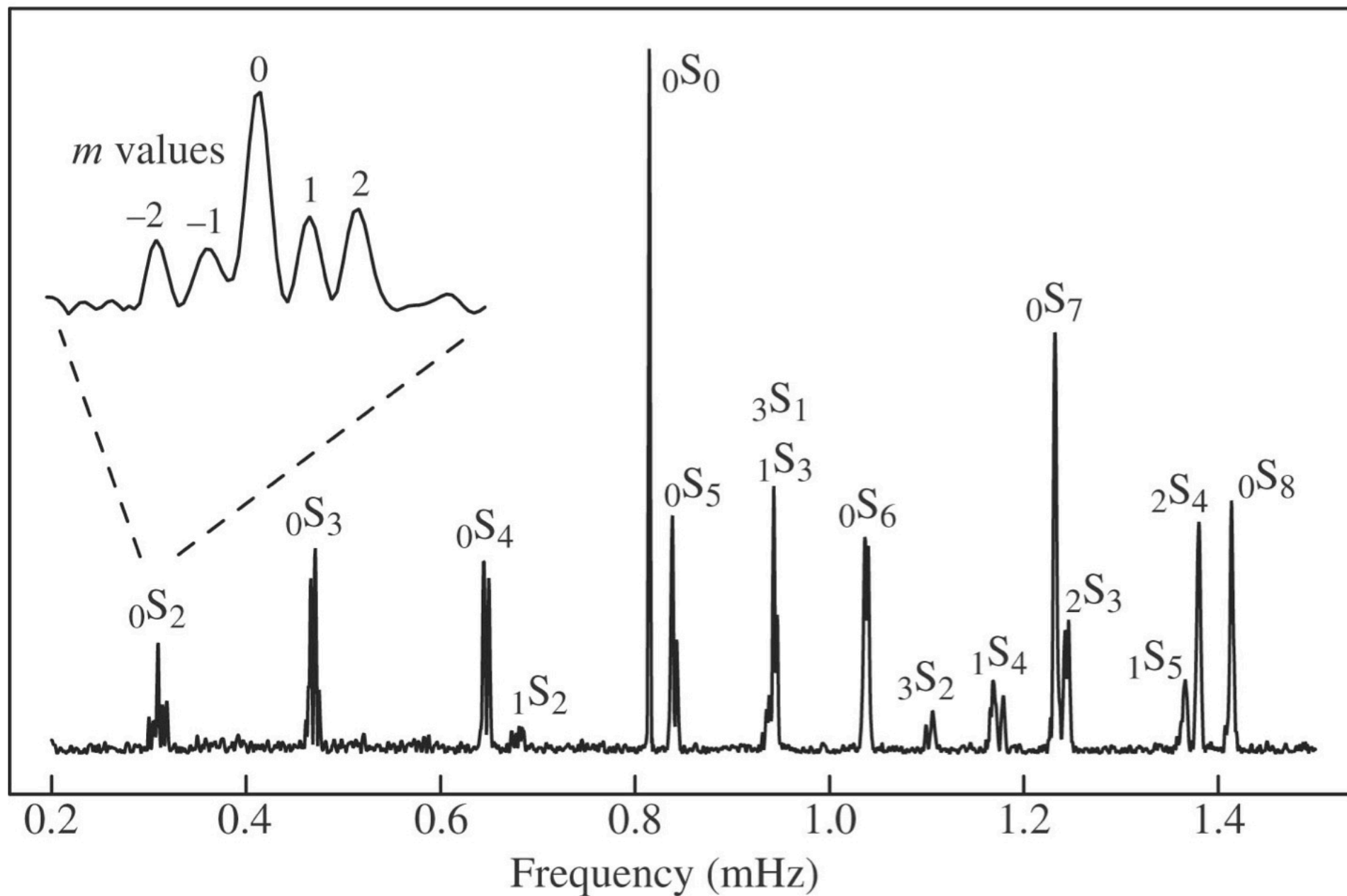
$2l+1$ singlets have different frequency



splitting and coupling



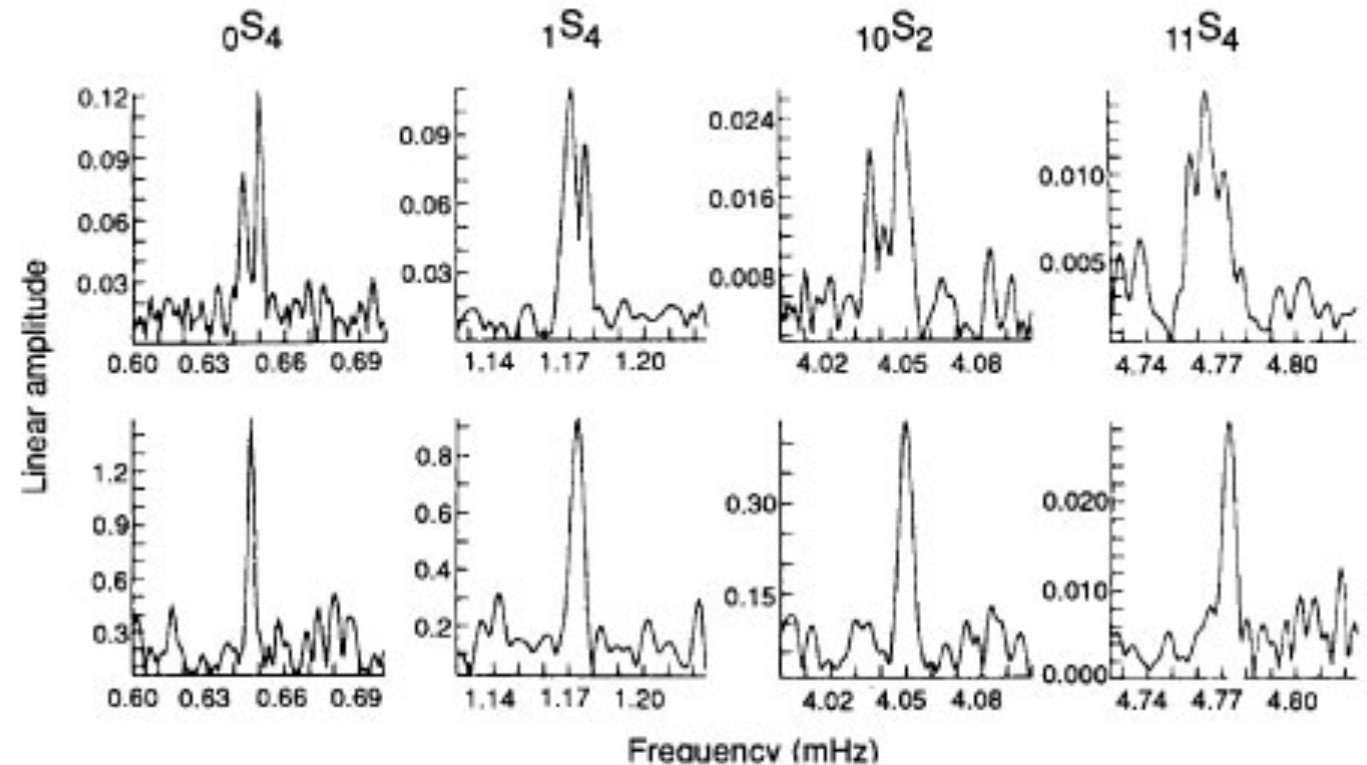
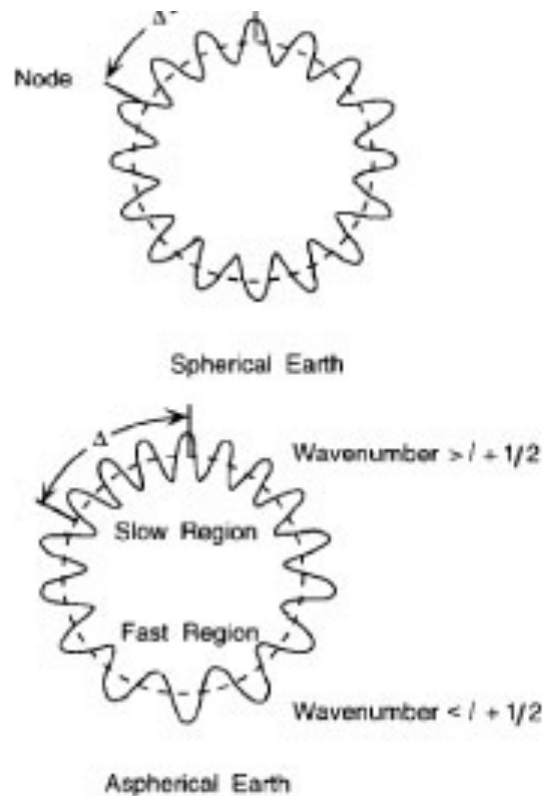
Singlet splitting



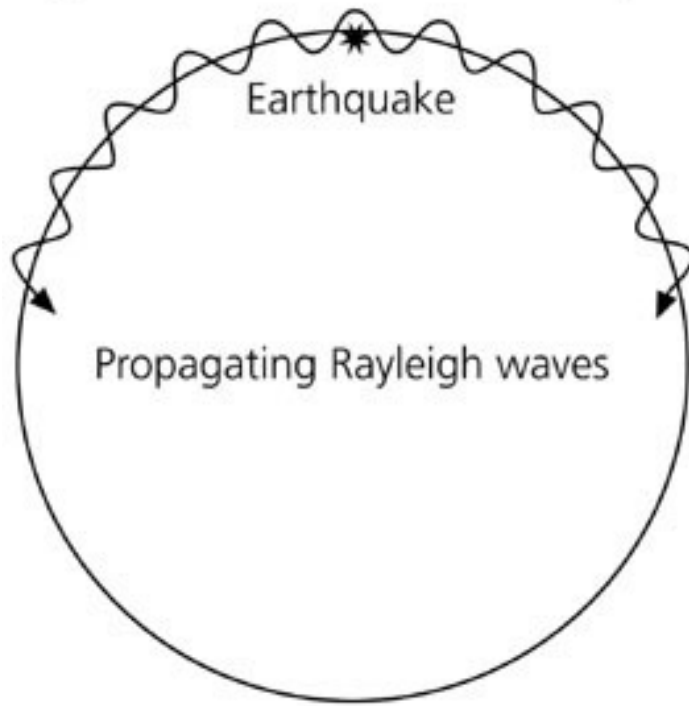
Splitting

Split due
to rotation

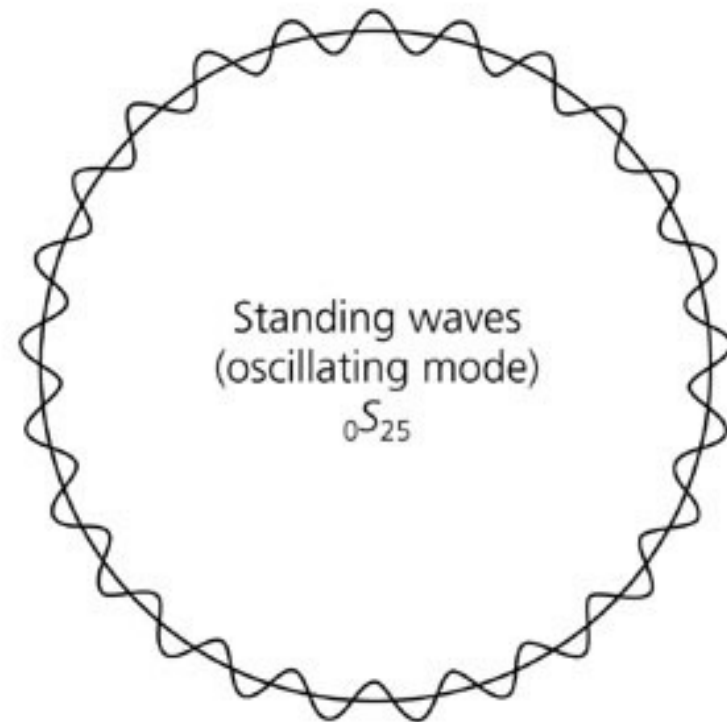
Anomalous split due
to inner core anisotropy



Normal modes and surface waves



A few minutes after
the earthquake



A few hours after
the earthquake

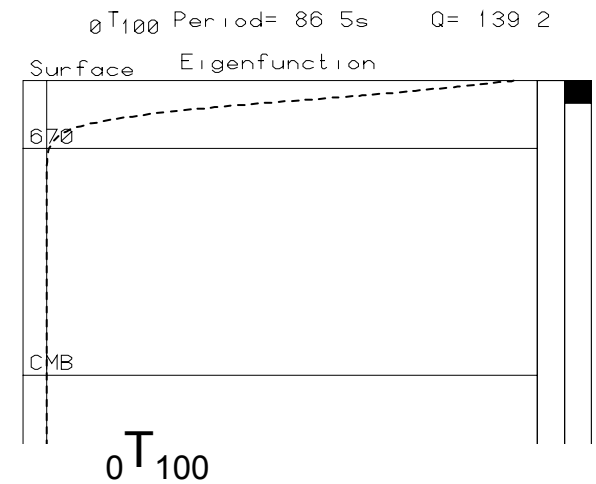
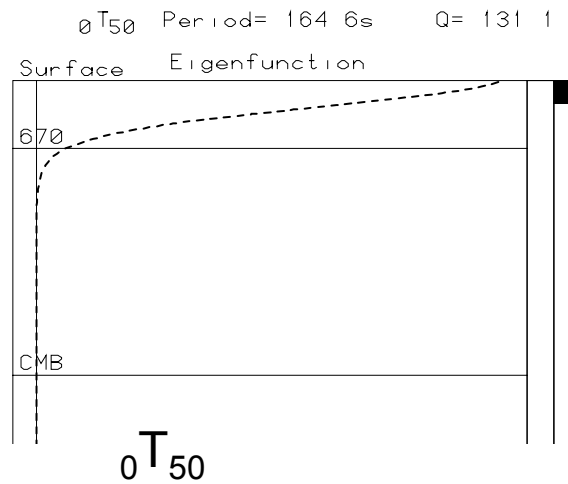
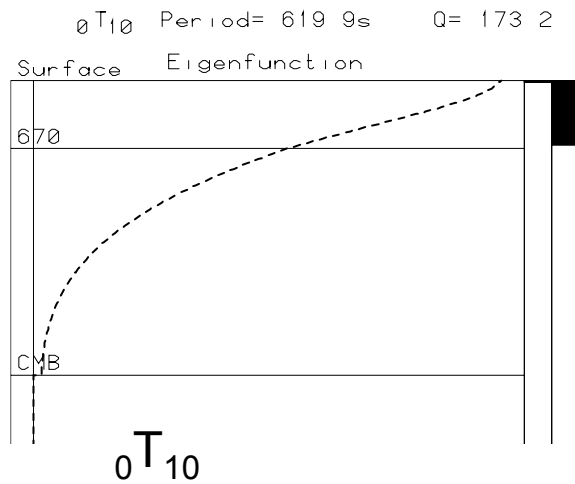
The mode with angular order l and frequency ${}_n\omega_l$ corresponds to a travelling wave with horizontal wavelength $\lambda_x = 2\pi / |k_x| = 2\pi a / (l + 1/2)$ that has $l+1/2$ wavelengths around the Earth.

These waves travel at horizontal phase velocity

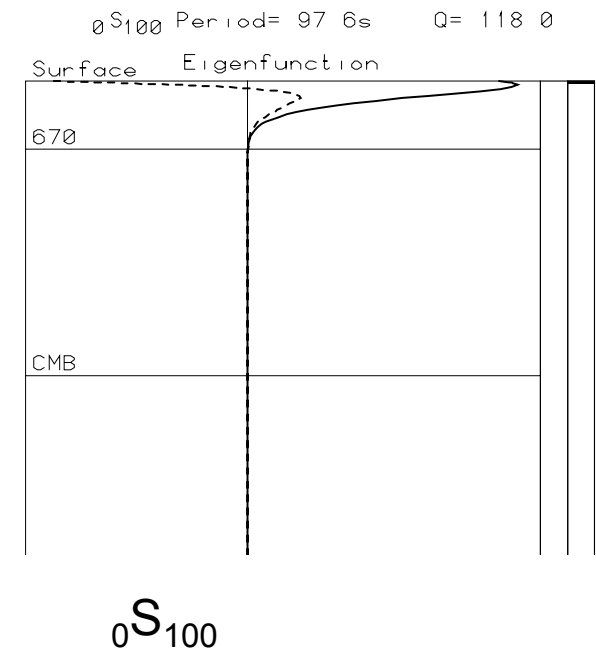
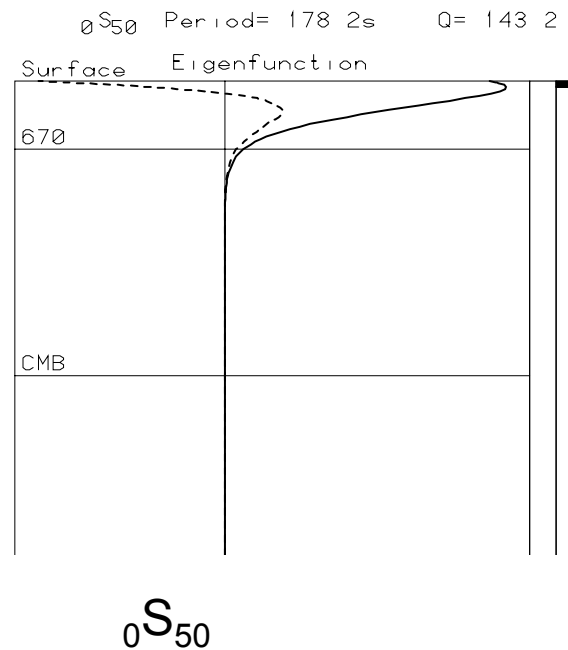
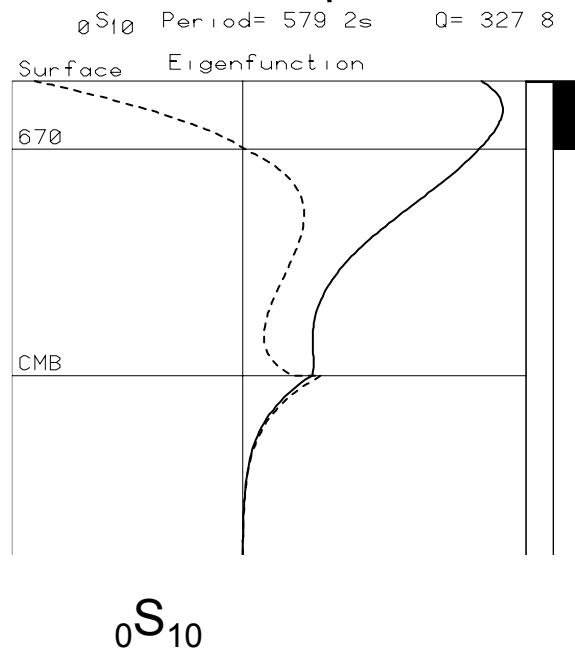
$$c_x = {}_n\omega_l / |k_x| = {}_n\omega_l a / (l + 1/2)$$

Eigenfunctions of fundamental modes

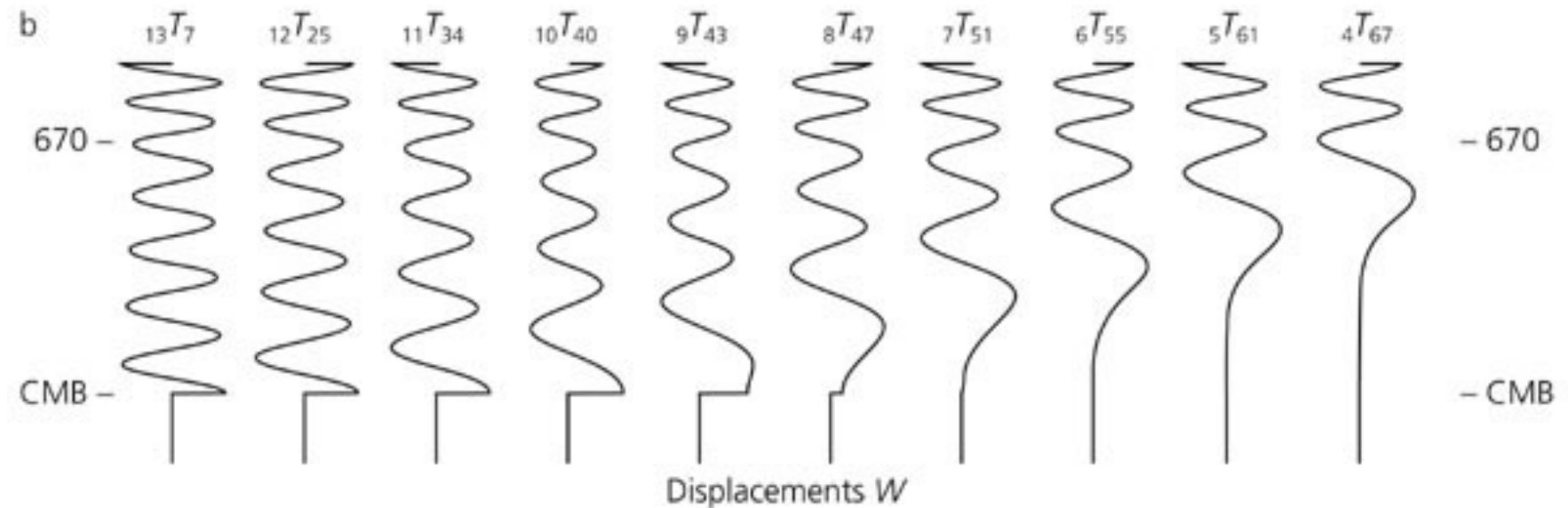
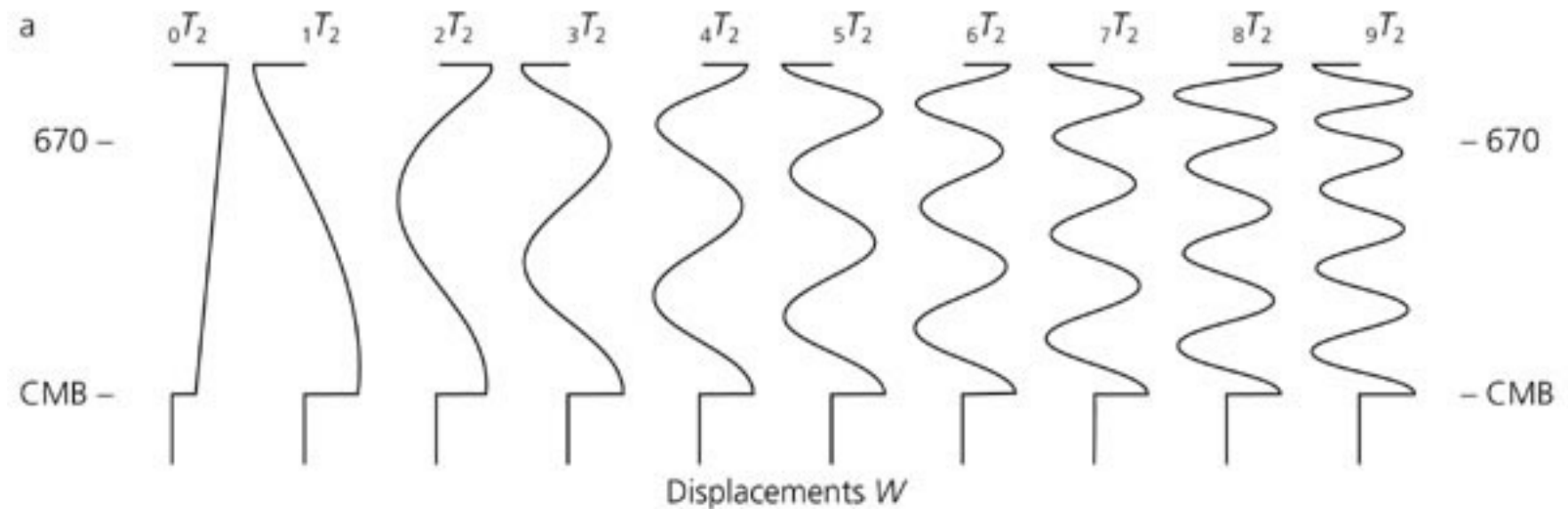
Fundamental toroidal branch



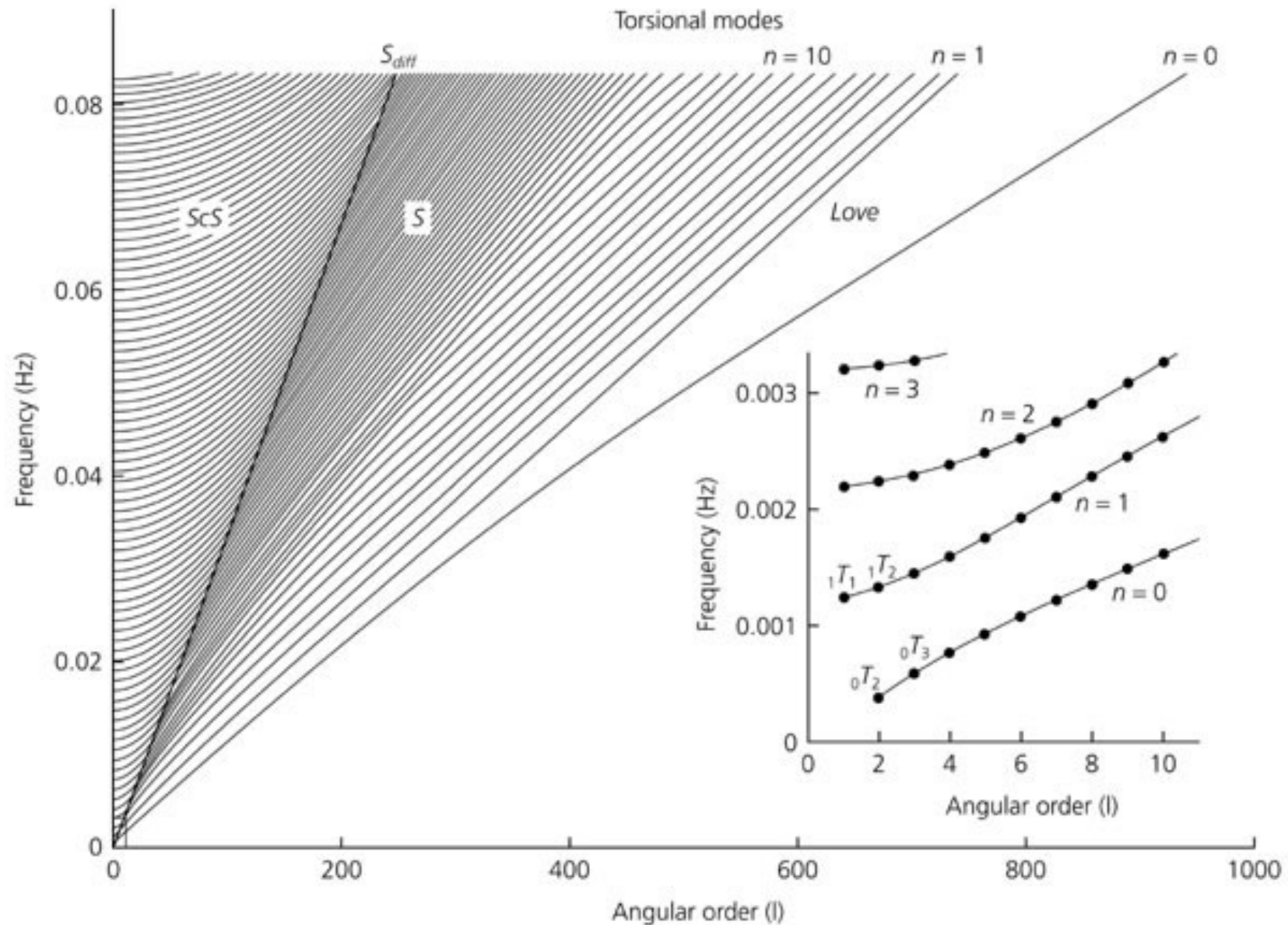
Fundamental spheroidal branch



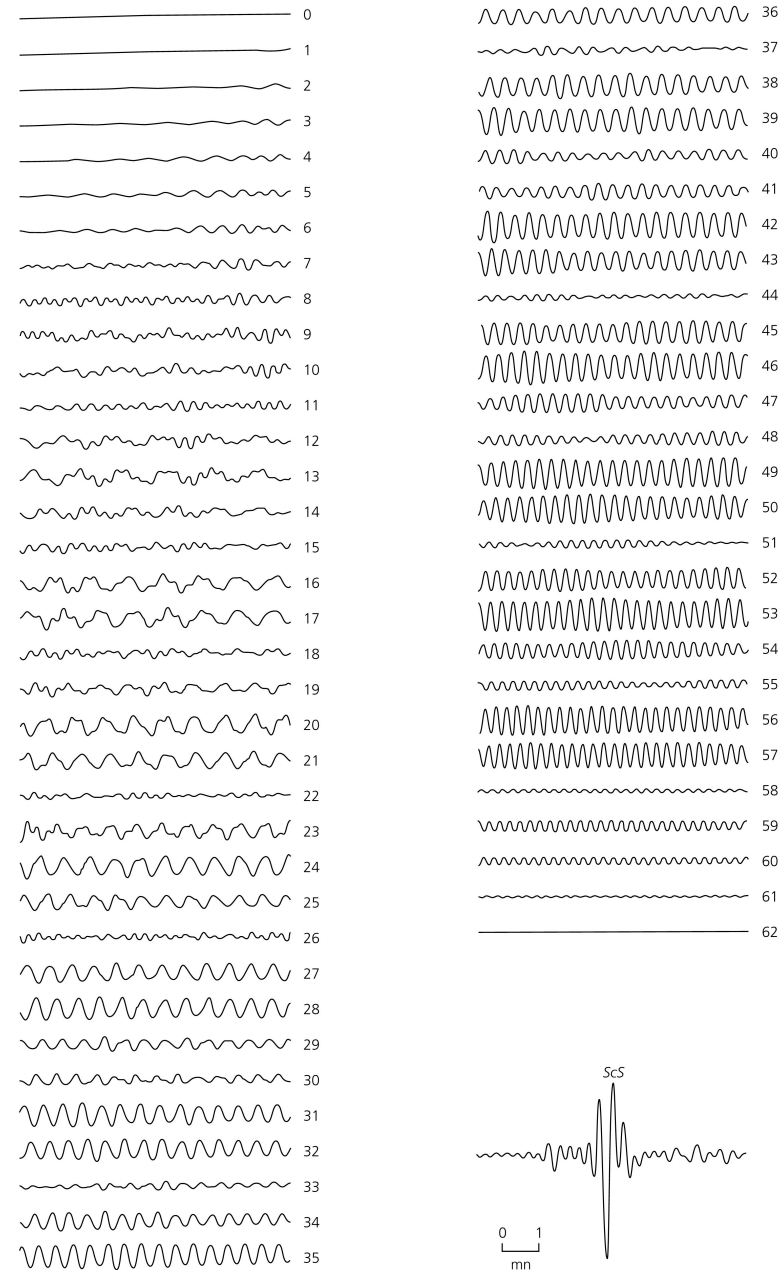
Eigenfunctions of overtones (higher modes)



Body waves by mode summation



ScS wave by mode summation



Application: density tomography

Only normal modes are sensitive to density perturbations in the Earth.

